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MATHEMATICAL ANALYSIS OF A DUPUIT-RICHARDS MODEL

SAFAA AL NAZER, CAROLE ROSIER, MUNKHGEREL TSEGMID

ABSTRACT. This article is concerned with a model which is an alternative to the 3D-Richards equation to describe the flow of water in shallow aquifers. The model couples the two dominant types of flow existing in the aquifer. The first is described by the classic Richards problem in the upper capillary fringe. The second results from Dupuit's approximation after vertical integration of the conservation laws between the bottom of the aquifer and the saturation interface. The final model consists of a strongly coupled system of parabolic-type pde which are defined in a time-dependent domain. First, we show how taking the low compressibility of the fluid into account eliminates the nonlinearity in the time derivative of the Richards equation. Then, the general framework of parabolic equations is used in non-cylindrical domains to give a global in time existence result to this problem.

1. INTRODUCTION

Populated areas are increasingly affected by contamination of soil and groundwater. Many modeling approaches are developed to study the vulnerability of aquifers to agricultural pollution, with a particular focus on the supply of nitrates. There is a wide variety of involved processes (chemical, hydrogeological, anthropic, ...) acting in a wide range of temporal and geometrical length scales. But we can notice that the main point for the derivation of the hydrogeological model is linked to a good description of the flow between the ground level (the level of the anthropic processes) and the water table. This will be crucial when studying the transport of chemical components in the aquifer. Indeed, it turns out that many chemical reactions are expected in the first meters of the subsoil, where oxygen is still very present. In particular, chemical species that reach the water table are not necessarily the same as those that have left the surface. This yields different speeds of the reactive kinetics. As a result, for an efficient mathematical modeling, the time upscaling process in this zone must keep track of all the time scales.

Only the hydrogeological question will be considered in this work. Aquifers are often characterized by a form of stratification of flows which enables the definition of interfaces. The slowness of the natural dynamics ensures a smooth and stable behavior for the interfaces. Besides, due to the dimensions of the aquifer, the flow

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can be assumed essentially orthogonal to the equipotential (Dupuit's hypothesis). The vertical integration of the Richards equation is thus possible, at least in the saturated zone. In this spirit, many 2D models have been developed and used since the 1960s (see for example the works of Jacob Bear, [5, 6]). For more historical notes on the origin of groundwater modeling, we refer interested readers to [14, 15, 17, 24]. But the approach by vertical integration is only valuable for very precise length and time scales, the time scale in particular being completely different from the typical durations of chemical reactions. However, such 2D models are widely used, although it is particularly difficult to correctly couple them to the flow in the unsaturated part of the basement. Several numerical studies have been conducted in this direction. Let us mention the work of [20] where the integrated model is directly coupled with a surface model. In [7] and [30], the coupling of the surface and underground flows is done using a Richards equation associated with a Signorini boundary condition (for the surface behavior). A class of models is proposed in [8], which consists in coupling purely vertical models (to describe the flow at a small time scale) with an horizontal model (describing the flow at a long time scale). They admit the same behavior than the 3D-Richards model for any time scale when the aquifer presents a small deepness compared to its large horizontal dimensions. They describe the essentially horizontal flow of a water table and the essentially vertical water supply flux from the surface through the unsaturated part between the groundwater and the ground level. In [27] a presentation of a rather similar model can be found, coupling 1D-Richards equation with a simplified model in the saturated part. Finally, in [1], this kind of model is integrated into a computational code called "SHE" (for "European Hydrological System" and later became SHETRAN), in the case where the water table remains away from ground level.

In this paper, a model belonging to the "Dupuit-Richards" model class is analyzed. Indeed, the 3D-Richards equation is considered in the capillary fringe while a vertical average of the mass conservation law is made in the saturated zone of the aquifer. Pressure and normal fluxes transmission properties are imposed at the saturation interface.

This model differs slightly from the one described in [8], firstly because the complete 3D Richards equations are considered in the unsaturated part and not only the vertical component of the flow. Nevertheless, the main difference lays in the consideration of the low compressibility of the fluid. This makes it possible to perform a change of variable which eliminates the nonlinearity of the time derivative of the moisture content, but also to treat the degeneracy in the equation governing the horizontal flow in the saturated part of the aquifer. On the other hand, the coupling of flows between the two areas of the aquifer results from the continuity property of the normal component of the flux at the interface of saturation, ensuring the mass conservation at the interface.

Of course, the numerical behavior of this model should be similar to the one obtained in [8], especially when the horizontal component of the hydraulic conductivity tends towards zero in the unsaturated part of the aquifer. Finally, the coupling of the 1D Richards equation with the 2D Dupuit approximation numerically justifies this model since it is less expensive than the complete resolution of the 3D Richards equation in terms of CPU time.

The mathematical study of the model is particularly delicate because of nonlinearities, the free boundary between each area of the aquifer and the difficulty resulting from the coupling between the two zones, which is expressed here in terms of flux at the interface. Finally, there is a general mathematical complexity in the structure of the set of PDEs modeling the dynamics of underground water. Indeed, when considering a free water table, we must face the gradual disappearance of water in the desaturation zone and thus the disappearance of a main unknown of the problem hence the mathematical challenges inherent in Richards equations. There exists a huge literature regarding the classical Richards equations. Let us mention the works of Alt *et al* ([3, 4]) and the papers [10, 18, 28] devoted to the study of the degenerate in time equation

$$\partial_t \theta(p) - \Delta p = 0,$$

where $\theta(p)$ denotes the moisture content. In the one-dimensional case, we also quote the work of Yin ([33]) concerning the existence of weak solution for the fully degenerate problem

$$\partial_t \theta(p) - \partial_x (\kappa(\theta(p)) \partial_x p) = 0,$$

when just assuming that $\theta', \kappa' > 0$.

Classically, the Kirchoff transform is applied to the Richards equation (under appropriate assumptions about porosity and permeability) to eliminate the nonlinearity in the diffusive term. In this work, the hypothesis of low compressibility of water is exploited instead, to eliminate the nonlinearity in the time derivative of the Richards equation. Even if this term is assumed to be very low, its very existence allows to define the one-to-one transformation that absorbs the degeneracy of the moisture content.

This transformation brings us back to the framework of quasilinear parabolic equations on non-cylindrical domains to which the auxiliary domain method introduced by Lions and Mignot [22, 25] can be applied to deal with the free boundary. Regarding the flow in the saturation zone, the vertical mean of conservation laws in this part leads to a degenerate elliptic equation whose degeneracy depends on the thickness of the zone. In addition, taking the compressibility of water into account introduces degeneracy (in the time derivative) also depending on the thickness of the saturated zone. A change of variable then allows to absorb the two degenerated terms and return to a regular parabolic equation.

The document is organized as follows: In section 2, the model coupling 3D Richards equation with the Dupuit horizontal approximation is introduced; consequences taking the compressibility of the fluid into account in the modeling are particularly detailed. Then a global in time existence result is given in Section 3 as well as preliminary results about the auxiliary domains method. The proof of the Theorem is performed in section 4. It consists in a fixed point strategy to deal with the difficulties associated with nonlinearities and coupling.

2. DERIVATION OF THE MODEL

The basis of the modeling is the mass conservation law written for fresh water coupled with the classical Darcy law for porous media. First, the Richards equations are obtained by taking the water compressibility into account. Then, the case of the unsaturated zone is distinguished from that of the saturated zone as explained in

the introduction. For the three-dimensional description, we denote by $\mathbf{x} := (x, z)$, $x = (x_1, x_2) \in \mathbb{R}^2$, $z \in \mathbb{R}$, the usual coordinates.

2.1. Geometry. The aquifer is represented by a three-dimensional domain $\Omega := \Omega_x \times (h_{\text{bot}}, h_{\text{soil}})$, $\Omega_x \subset \mathbb{R}^n$ with $n \geq 2$ ($x = (x_1, x_2)$), function h_{bot} (respect. h_{soil}) describing its lower (respect. upper) topography. The upper and lower surfaces are thus defined by the graph of the functions $h_{\text{bot}} = h_{\text{bot}}(x)$ and $h_{\text{soil}} = h_{\text{soil}}(x)$, $x \in \Omega_x$. We assume that

$$h_{\text{soil}}(x) > h_{\text{bot}}(x), \quad \forall x \in \Omega_x. \quad (2.1)$$

More precisely the domain is given by:

$$\Omega = \left\{ (x, z) \in \Omega_x \times \mathbb{R} \mid z \in]h_{\text{bot}}(x), h_{\text{soil}}(x)[\right\}. \quad (2.2)$$

We always denote by $\vec{\nu}$ the outward unit normal and \vec{e}_3 is the unitary vertical vector pointing up. The boundary $\partial\Omega$ of Ω is divided into three zones (bottom, top and vertical)

$$\partial\Omega = \Gamma_{\text{bot}} \sqcup \Gamma_{\text{soil}} \sqcup \Gamma_{\text{ver}},$$

with

$$\begin{aligned} \Gamma_{\text{bot}} &:= \left\{ (x, z) \in \Omega \mid z = h_{\text{bot}}(x) \right\}, & \Gamma_{\text{soil}} &:= \left\{ (x, z) \in \Omega \mid z = h_{\text{soil}}(x) \right\}, \\ \Gamma_{\text{ver}} &:= \left\{ (x, z) \in \Omega \mid x \in \partial\Omega_x \right\}. \end{aligned}$$

The model divides the flow description into two subregions of Ω (possibly time-dependent) in each of which the flow presents a different behavior. We denote by h the depth of the free interface separating the freshwater layer and the unsaturated part of the aquifer. The definition of these zones is thus based on the function $h = h(t, x)$ which is an unknown of our problem. Then, for a given function $h = h(t, x)$ such that $h_{\text{bot}} \leq h \leq h_{\text{soil}}$, we introduce:

$$\Omega_t^- := \left\{ (x, z) \in \Omega \mid z < h(t, x) \right\} \quad \text{and} \quad \Omega_t := \left\{ (x, z) \in \Omega \mid z > h(t, x) \right\}, \quad (2.3)$$

and

$$\Gamma_t := \left\{ (x, z) \in \Omega \mid z = h(t, x) \right\}. \quad (2.4)$$

2.2. Conservation laws. In view of the (large) dimensions of an aquifer (related to the characteristic size of the porous structure of the underground), we consider a continuous description of the porous medium. The effective velocity q of the flow is thus related to the pressure P through the Darcy law associated with a nonlinear anisotropic conductivity

$$q = -\frac{\kappa(P)K_0}{\mu}(\nabla P + \rho g \nabla z),$$

where ρ and μ are respectively the density and the viscosity of the fluid, K_0 is the permeability of the soil, $\kappa(P)$ is the relative conductivity and g the gravitational acceleration constant. Introducing the hydraulic head H defined by

$$H = \frac{P}{\rho_0 g} + z, \quad (2.5)$$

the previous equation is written as follows:

$$q = -K \nabla H - \frac{\kappa(P)K_0}{\mu}(\rho - \rho_0)g \nabla z, \quad K = \frac{\kappa(P)K_0 \rho_0 g}{\mu}. \quad (2.6)$$

In this relation, the matrix K is the hydraulic conductivity which expresses the ability of the underground to conduct the fluid. ρ_0 denotes the reference density of the fluid. Next, the conservation of mass during displacement is given by the following equation

$$\partial_t(\theta\rho) + \nabla \cdot (\rho q) = \rho Q, \quad (2.7)$$

where Q denotes a generic source term (for production and replenishment).

The function θ is the volumetric moisture content defined by

$$\theta = \phi s,$$

where ϕ is the porosity of the medium and s is the saturation. If the air present in the unsaturated zone is assumed to have infinite mobility, the saturation s , and then the function θ are thus considered monotone functions depending on the pressure as we will detail latter.

2.3. State equation for the fluid compressibility. The fluid is considered compressible by assuming that pressure P is related to the density ρ as follows (cf. [11]):

$$\frac{d\rho}{\rho} = \alpha_P dP \Leftrightarrow \rho = \rho_0 e^{\alpha_P(P-P_0)}. \quad (2.8)$$

The real number $\alpha_P \geq 0$ represents the fluid compressibility coefficient and P_0 is the pressure of reference. Further assuming $\alpha_P = 0$ we would recover the incompressible case.

This compressibility coefficient α_P is of course very low but, as mentioned in the introduction, the fact of not completely neglecting it makes it possible to facilitate the mathematical analysis of the Richards equations.

2.4. Permeability tensor K_0 . The nonlinear hydraulic conductivity K is given by $K = \frac{\kappa(P)\rho_0 g}{\mu} K_0$. The soil transmission properties are characterized by the porosity function ϕ and the permeability tensor $K_0(x, z)$. The matrix K_0 is a 3×3 symmetric positive definite tensor which describes the conductivity of the *saturated* soil at the position $(x, z) \in \Omega$. We introduce $K_{xx} \in \mathcal{M}_{22}(\mathbb{R})$, $K_{zz} \in \mathbb{R}^*$ and $K_{xz} \in \mathcal{M}_{21}(\mathbb{R})$ such that

$$K_0 = \begin{pmatrix} K_{xx} & K_{xz} \\ K_{xz}^T & K_{zz} \end{pmatrix}. \quad (2.9)$$

2.5. Hypothesis. Let us now list the assumptions on the fluid and medium characteristics but also on the flow which are meaningful in the context of this problem.

2.5.1. Hypothesis on the fluid and on the medium. In the model, the effects of the rock compressibility are neglected, the porosity of the medium ϕ do not depend on the pressure variations and it is thus assumed to be a constant.

Compressibility of the fluid. The fluid (namely here fresh water) is weakly compressible. It means that

$$\alpha_P \ll 1. \quad (2.10)$$

Let us exploit this assumption. In natural conditions and especially in an aquifer, one observes small fluid mobility (defined by the ratio κ/μ). First consequence of the low compressibility of the fluid combined with the low mobility of fluid appears in the momentum equation. We perform a Taylor expansion with regard to P of the

density ρ in the gravity term of the Darcy equation. Neglecting the terms weighted by $\alpha_P \kappa / \mu \ll 1$ in (2.6), we get:

$$q = -K \nabla H, \quad K = \frac{\kappa(P) \rho_0 g}{\mu} K_0. \quad (2.11)$$

Second consequence is $\nabla \rho \cdot q \ll 1$ which leads to the following simplification in the mass conservation equation (2.7):

$$\rho \partial_t \theta + \theta \partial_t \rho + \rho \nabla \cdot q = \rho Q.$$

Neglecting in this way the variation of density in the direction of flow is sometimes considered as an extra assumption called Bear's hypothesis (*cf* [2]). Including (2.8), that is $\partial_t \rho = \rho \alpha_P \partial_t P$ in the latter equation, we get

$$\rho \partial_t \theta + \rho \theta \alpha_P \partial_t P + \rho \nabla \cdot q = \rho Q.$$

After simplification by $\rho > 0$, we finally obtain

$$\partial_t \theta + \theta \alpha_P \partial_t P + \nabla \cdot q = Q. \quad (2.12)$$

Equivalently, using the hydraulic head (2.5) and the Darcy law (2.11), (2.12) can be written

$$\partial_t \theta + S_0 \partial_t H - \nabla \cdot (K \nabla H) = Q \quad \text{where} \quad S_0 = \rho_0 g \phi \alpha_P. \quad (2.13)$$

We notice that if the fluid is assumed incompressible, $\alpha_P = 0$, then Eq. (2.12) is the classical Richards equation in pressure formulation. An adequate definition of the volumetric moisture content θ and of the mobility function κ is the key of the model.

Richards hypothesis. The Richards model is moreover based on the assumption that the air pressure in the underground equals the atmospheric pressure, thus is not an unknown of the problem. One thus assumes that the saturation and the relative conductivity of the soil are given as *functions* of the fluid pressure P , denoted respectively by $s = s(P)$ and $\kappa = \kappa(P)$. We introduce the saturation pressure P_s which is a fixed real number. The fully-saturated part of the medium corresponds to the region $\{\mathbf{x}; P(\cdot, \mathbf{x}) > P_s\}$, while it is partially-saturated in the capillary fringe $\{\mathbf{x}; P_d < P(\cdot, \mathbf{x}) \leq P_s\}$. The dry part is defined by the set $\{\mathbf{x}; P(\cdot, \mathbf{x}) \leq P_d\}$. The moisture content is such that

$$\theta = \begin{cases} \phi & \text{(saturated zone)} & \text{if } P(\cdot, \mathbf{x}) > P_s, \\ \theta(P) & \text{(with } 0 \leq \theta(P) \leq \phi \text{ and } \theta'(P) > 0) & \text{if } P_d < P(\cdot, \mathbf{x}) \leq P_s, \\ \theta_0 = \phi s_0 & \text{(dry zone)} & \text{if } P(\cdot, \mathbf{x}) \leq P_d, \end{cases} \quad (2.14)$$

where $s_0 > 0$ corresponds to a residual saturation which is positive. The associated relative hydraulic mobility is then defined by

$$\kappa(P) = \begin{cases} 1 & \text{(saturated zone)} & \text{if } P(\cdot, \mathbf{x}) > P_s, \\ \kappa(\theta(P)) & \text{(with } 0 \leq \kappa(P) \leq 1 \text{ and } (\kappa \circ \theta)'(P) > 0) & \text{if } P_d < P(\cdot, \mathbf{x}) \leq P_s, \\ 0 & \text{(dry zone)} & \text{if } P(\cdot, \mathbf{x}) \leq P_d. \end{cases} \quad (2.15)$$

There is a large choice of available models for s and κ . The most classical examples for an air-water system are the van Genuchten model [32] with no-explicit dependence on the bubbling pressure but with fitting parameters, and the Brooks and Corey model [9]. The important point is that these models are such that

$$s(P) = 1 \iff P \geq P_s \quad \kappa(P) = 1 \iff P \geq P_s. \quad (2.16)$$

In particular, the water pressure is greater than the bubbling pressure P_s if and only if the soil is completely saturated .

2.5.2. Hypothesis on the flow. The following assumption is introduced for upscaling the 3D problem to a 2D model in the saturated part of the domain.

Dupuit approximation (hydrostatic approach) Dupuit assumption consists in considering that the hydraulic head is constant along each vertical direction (vertical equipotentials). It is legitimate since one actually observes quasi-horizontal displacements when the thickness of the aquifer is small compared to its width and its length and when the flow is far from sinks and wells.

2.6. Model coupling vertical 3d-Richards flow and Dupuit horizontal flow. From the compressible Richards equation (2.12), we will deduce the equations governing the flow in each zone of the aquifer.

• **Three-dimensional Richards equation in the upper capillary fringe.**

In the unsaturated part of the aquifer, Ω_t , the 3D Richards equation (2.12) holds

$$\begin{cases} \partial_t \theta + \theta_{\alpha P} \partial_t P + \nabla \cdot q = Q & \text{for } (t, x, z) \in (0, T) \times \Omega_t, \\ q \cdot \vec{\nu} = 0 & \text{for } (t, x, z) \in (0, T) \times (\Gamma_{\text{soil}} \cup \Gamma_{\text{ver}}), \\ P(t, x, h(t, x)) = P_s & \text{for } (t, x) \in (0, T) \times \Omega_x, \\ P(0, x, z) = P_{\text{init}}(x, z) & \text{for } (x, z) \in \Omega_0. \end{cases} \quad (2.17)$$

The effective velocity q is given by

$$q = -K \nabla \left(\frac{P}{\rho_o g} + z \right), \quad K = \frac{\kappa(P) K_0 \rho_o g}{\mu}.$$

We emphasize that the model (2.17) depends by definition on the depth h which is expected to belong to the interval $(h_{\text{bot}}, h_{\text{soil}})$.

• **Dupuit horizontal flow in the saturated zone.**

In the saturated part of the aquifer, Ω_t^- , the vertical average of the 3D Richards equation (2.13) describes the horizontal flow of this part, thus reducing the 3D problem to a 2D problem. For the upscaling procedure, we proceed as it was done in the context of the seawater intrusion in [11, 12].

Upscaling procedure

The vertical integration is performed between depths h_{bot} and h . Since $\theta(P) = \phi$ in the saturated zone, the vertical average (2.13) leads to

$$\int_{h_{\text{bot}}}^h (S_0 \partial_t H + \nabla \cdot q) dz = \int_{h_{\text{bot}}}^h Q dz.$$

$B_f = h - h_{\text{bot}}$ denotes the thickness of the saturated zone and \tilde{Q} the source term representing distributed surface supply of fresh water into the free aquifer:

$$\tilde{Q} = \frac{1}{B_f} \int_{h_{\text{bot}}}^h Q dz.$$

Applying Leibnitz rule to the first term in the left-hand side yields:

$$\int_{h_{\text{bot}}}^h S_0 \partial_t H dz = S_0 \frac{\partial}{\partial t} \int_{h_{\text{bot}}}^h H dz - S_0 H|_{z=h} \partial_t h + S_0 H|_{z=h_{\text{bot}}} \partial_t h_{\text{bot}}.$$

We denote by \tilde{H} the vertically averaged hydraulic head

$$\tilde{H} = \frac{1}{B_f} \int_{h_{bot}}^h H dz.$$

Because of Dupuit approximation, $H(x_1, x_2, z) \simeq \tilde{H}(x_1, x_2)$, $x = (x_1, x_2) \in \Omega$, $z \in (h_{bot}, h)$, hence we have

$$\int_{h_{bot}}^h S_0 \partial_t H dz = S_0 B_f \partial_t \tilde{H}.$$

Besides

$$\int_{h_{bot}}^h \nabla \cdot q dz = \nabla' \cdot (B_f \tilde{q}') + q_{|z=h^-} \cdot \nabla(z-h) - q_{|z=h_{bot}^+} \cdot \nabla(z-h_{bot}),$$

where $\nabla' = (\partial_{x_1}, \partial_{x_2})$, $q' = (q_{x_1}, q_{x_2})$. The averaged Darcy velocity

$\tilde{q}' = \frac{1}{B_f} \int_{h_{bot}}^h q' dz$ is given by

$$\tilde{q}' = -\frac{1}{B_f} \int_{h_{bot}}^h (K \nabla' H) dz = -\tilde{K} \nabla' \tilde{H}, \quad \tilde{K} = \frac{1}{B_f} \int_{h_{bot}}^h \frac{K_0 \rho_0 g}{\mu} dz,$$

(we remind that $\kappa(P) = 1$ for $z \in (h_{bot}, h)$). The averaged mass conservation law for the freshwater in the saturated zone finally reads

$$\begin{aligned} S_0 B_f \partial_t \tilde{H} &= \nabla' \cdot (B_f \tilde{K} \nabla' \tilde{H}) \\ &+ q_{|z=h_{bot}^+} \cdot \nabla(z-h_{bot}) - q_{|z=h^-} \cdot \nabla(z-h) + B_f \tilde{Q}. \end{aligned} \quad (2.18)$$

In this equation, the term $B_f \tilde{K}$ may be viewed as the dynamic transmissivity of the freshwater layer. At this stage, we have obtained an undetermined system of two pdes ((2.17)-(2.18)) with three unknowns P , \tilde{H} and h .

Fluxes and continuity equations across the interface

Our aim is now to include in the model the continuity and transfert properties across interfaces. As a consequence, we express the two flux terms appearing in (2.18) and the number of unknowns is reduced.

▷ Flux across the saturation interface:

The saturation interface is characterized by the cartesian equation

$F(x_1, x_2, z, t) = 0 \Leftrightarrow z - h(x_1, x_2, t) = 0$, so the unit normal vector $\vec{\nu}$ to the interface is colinear to $\nabla(z-h)$.

The relation ruling continuity of the normal component of the velocity thus reads

$$(q_{|z=h^+} - q_{|z=h^-}) \cdot \vec{\nu} = 0 \Leftrightarrow q_{|z=h^+} \cdot \nabla(z-h) = q_{|z=h^-} \cdot \nabla(z-h). \quad (2.19)$$

▷ Approximation of the flux $q_{|z=h^+} \cdot \nabla(z-h)$:

The flux $q_{|z=h^+} \cdot \nabla(z-h)$ expresses mass transfers between the two parts of the aquifer. As it is done in [8], we approximate the flux by

$$q_{|z=h^+} \cdot \nabla(z-h) \simeq \int_{h(t,x)}^{h_{soil}(x)} (\phi \frac{\partial s(P)}{\partial t} + \phi s(P) \alpha_P \frac{\partial P}{\partial t} - Q) dz. \quad (2.20)$$

This approximation comes from the hypothesis of an almost null horizontal hydraulic conductivity (i.e. $K_{xx} \simeq (0)$) in the capillary fringe. This corresponds to a flow almost vertical in this part of the aquifer. So the 3D-Richards equation is reduced to an 1D-equation. Integrating this 1D equation between h and h_{soil} leads to the approximation (2.20).

It is an essential difference with the mathematical analysis presented in [31] in which the exchanges between the two parts of the aquifer were simplified and represented by the addition of an external source term, thus decoupling the two problems.

▷ Impermeable layer at $z = h_{soil}$:

Since the lower layer is impermeable, there is no flux across the boundary $z = h_{bot}$:

$$q(h_{bot}) \cdot \nabla(z - h_{bot}) = 0. \tag{2.21}$$

▷ Continuity equations:

Continuity relation imposed on the interface enables to properly reduce the number of unknowns in equations (2.17)-(2.18).

Dupuit approximation reads $\tilde{H} \simeq H|_{z=h^-}$, the pressure P thus satisfies in Ω_t^-

$$P(t, x, z) = \rho_0 g(\tilde{H}(t, x) - z) \quad \text{for } t \in [0, T[, \quad (x, z) \in \Omega_t^-. \tag{2.22}$$

Besides, the pressure is continuous across Γ_t , it follows that

$$P(t, x, h^-) = P(t, x, h^+) = P_s \Leftrightarrow \tilde{H} = \frac{P_s}{\rho_0 g} + h. \tag{2.23}$$

Equation (2.23) allows to substitute \tilde{H} by h in Eq. (2.18), so we have

$$S_0 B_f \partial_t h - \nabla' \cdot (B_f \tilde{K} \nabla' h) = B_f \tilde{Q} \tag{2.24}$$

$$- \int_{h(t,x)}^{h_{soil}(x)} \left(\phi \frac{\partial s(P)}{\partial t} + \phi s(P) \alpha_P \frac{\partial P}{\partial t} - Q \right) dz \text{ in } (0, T) \times \Omega_x,$$

$$\tilde{K} \nabla' h \cdot \vec{\nu} = 0 \text{ on } (0, T) \times \partial\Omega_x, \tag{2.25}$$

with

$$B_f = (h - h_{bot}), \quad \tilde{K} = \frac{1}{B_f} \int_{h_{bot}}^h \frac{K_0 \rho_0 g}{\mu} dz \quad \text{and} \quad S_0 = \rho_0 g \phi \alpha_P. \tag{2.26}$$

The homogeneous Neumann condition on $\partial\Omega_x$ is assumed to simplify the presentation.

The final model (\mathcal{M}) coupling 3D-Richards flow and Dupuit horizontal flow is completed with initial and boundaries conditions. It thus consists in the following system:

- In Ω_t the following 3d-Richards equation holds

$$\begin{cases} \partial_t \theta(P) + \theta \alpha_P \partial_t P + \nabla \cdot q = Q & \text{in } (0, T) \times \Omega_t, \\ q \cdot \vec{\nu} = 0 & \text{on } (0, T) \times (\Gamma_{soil} \cup \Gamma_{ver}), \\ P(t, x, h(t, x)) = P_s & \text{in } (0, T) \times \Omega_x, \\ P(0, x, z) = P_0(x, z) & \text{in } \Omega_0, \end{cases}$$

where the saturation pressure P_s is assumed to be constant with respect to the time and to the space. The effective velocity q is given by

$$q = -K \nabla \left(\frac{P}{\rho_0 g} + z \right), \quad K = \frac{\kappa(P) K_0 \rho_0 g}{\mu}.$$

- In Ω_t^- the pressure P satisfies

$$P(t, x, z) = \rho_0 g \left(\frac{P_s}{\rho_0 g} + h - z \right) \quad \text{in } (0, T) \times \Omega_t^-.$$

- The depth of Γ_t, h , satisfies in Ω_x

$$\begin{cases} S_0 B_f \partial_t h - \nabla' \cdot (B_f \tilde{K} \nabla' h) = B_f \tilde{Q} - \int_{h(t,x)}^{h_{soil}(x)} \left(\phi \frac{\partial s(P)}{\partial t} + \phi s(P) \alpha_P \frac{\partial P}{\partial t} - Q \right) dz, \\ \tilde{K} \nabla' h \cdot \vec{\nu} = 0 & \text{on } (0, T) \times \partial \Omega_x, \\ h(0, x) = h_0(x) & \text{in } \Omega_x. \end{cases}$$

3. MATHEMATICAL SETTING AND MAIN RESULTS

The problem (\mathcal{M}) being a problem with free boundary, we are going to define the general framework of parabolic equation in non cylindrical domain, introduced by Lions and Mignot respectively in [22] and [25]. There are obviously many other techniques to deal with free boundary problems, we refer interested readers to the recent paper [19] and references therein.

3.1. Notations and auxiliary results. For any $T > 0$, let \mathcal{O}_T be the open domain of $\mathbb{R}^+ \times \Omega$ defined by

$$\mathcal{O}_T = \{(t, x, z) \in (0, T) \times \Omega, h(t, x) < z\}$$

where h is the position of the interface Γ_t . We set

$$\Omega_t = \{(x, z) \in \Omega, z \in]h(t, x), h_{soil}[\}, \quad \mathcal{O}_T^c = ((0, T) \times \Omega) \setminus \mathcal{O}_T,$$

$$\Gamma = \partial \mathcal{O}_T \text{ (boundary of } \mathcal{O}_T), \quad \Gamma' = \Gamma \setminus (\Omega_0 \cup \Omega_T) \text{ (lateral boundary of } \mathcal{O}_T).$$

We define

$$H^{0,1}(\mathcal{O}_T) = \{u \in L^2(\mathcal{O}_T) \mid D_x^p u \in L^2(\mathcal{O}_T) \text{ for } |p| \leq 1\},$$

where

$$D_x^p u = \{D_x^\alpha u \mid \alpha = (\alpha_1, \alpha_2, \alpha_2) \text{ with } |\alpha| = p\}.$$

It is an Hilbert space endowed with the norm

$$\|u\|_{H^{0,1}(\mathcal{O}_T)} = \left(\sum_{p \leq 1} \int_{\mathcal{O}_T} |D^p u| dx dt \right)^{1/2}.$$

$F(\mathcal{O}_T)$ denotes the closure in $H^{0,1}(\mathcal{O}_T)$ of functions of $\mathcal{D}(\bar{\mathcal{O}}_T)$ null in a neighborhood of Γ_t and $F'(\mathcal{O}_T)$ its topological dual. Besides, we introduce

$$\mathcal{B}(\mathcal{O}_T) = \{u \mid u \in F(\mathcal{O}_T), \frac{du}{dt} \in F'(\mathcal{O}_T)\},$$

endowed with the Hilbertian norms $\|\cdot\|_{\mathcal{B}(\mathcal{O}_T)}^2 = \|\cdot\|_{F(\mathcal{O}_T)}^2 + \|\partial_t \cdot\|_{F'(\mathcal{O}_T)}^2$.

Finally, $B_0(\mathcal{O}_T)$ (resp. $B_T(\mathcal{O}_T)$) is the closure in $\mathcal{B}(\mathcal{O}_T)$ of functions of $\mathcal{B}(\mathcal{O}_T)$ null in a neighborhood of $t = 0$ (resp. $t = T$).

We now state some auxiliary results proved in [22]

Lemma 3.1. *If \mathcal{O}_T is sufficiently regular, we thus have*

- 1. $H^{0,1}(\mathcal{O}_T) = L^2([0, T]; H^1(\Omega_t))$ where

$$L^2(0, T; H^1(\Omega_t)) = \{u \mid u(t, \cdot) \in H^1(\Omega_t), t \in [0, T] \text{ a.e. and } \|u\|_{H^{0,1}(\mathcal{O}_T)} < +\infty\},$$

$$\text{with } \|u\|_{H^{0,1}(\mathcal{O}_T)} = \int_0^T \|u\|_{H^1(\Omega_t)}^2 dt.$$

A similar result holds for $F(\mathcal{O}_T)$.

- 2. For $u \in F(\mathcal{O}_T)$, we can define $\gamma(u)$, the trace of u on Γ' in $L^2(\Gamma')$. Moreover $u \in F(\mathcal{O}_T) \iff \gamma(u) = 0$ on Γ_t .
- 3. Let $u \in \mathcal{B}(\mathcal{O}_T)$, thus $u \in B_T(\mathcal{O}_T) \iff u(T, \cdot) = 0$.

- 4. $\forall u, v \in \mathcal{B}(\mathcal{O}_s)$, we have

$$\left\langle \frac{\partial u}{\partial t}, v \right\rangle_{F',F} + \left\langle \frac{\partial v}{\partial t}, u \right\rangle_{F',F} = (u(s, \cdot), v(s, \cdot))_{L^2(\Omega_s)} - (u(0, \cdot), v(0, \cdot))_{L^2(\Omega_0)}. \quad (3.1)$$

For the sake of brevity, we shall write $H^1(\Omega) = W^{1,2}(\Omega)$ and

$$V(\Omega) = H_{0,\Gamma_{bot}}^1(\Omega) = \{u \in H^1(\Omega), u = 0 \text{ on } \Gamma_{bot}\}, V'(\Omega) = (H_{0,\Gamma_{bot}}^1(\Omega))'.$$

The embeddings $V(\Omega) \subset L^2(\Omega) \subset V'(\Omega)$ are dense and compact. For any $T > 0$, let $W_0(0, T, \Omega)$ denotes the space

$$W_0(0, T, \Omega) := \{\omega \in L^2(0, T; V(\Omega)), \partial_t \omega \in L^2(0, T; V'(\Omega))\},$$

endowed with the Hilbertian norm $\|\cdot\|_{W_0(0,T,\Omega)}^2 = \|\cdot\|_{L^2(0,T;V(\Omega))}^2 + \|\partial_t \cdot\|_{L^2(0,T;V'(\Omega))}^2$.

The following embeddings are continuous ([23] prop. 2.1 and thm 3.1, chapter 1)

$$W_0(0, T, \Omega) \subset \mathcal{C}([0, T]; [V(\Omega), V'(\Omega)]_{\frac{1}{2}}) = \mathcal{C}([0, T]; L^2(\Omega))$$

while the embedding

$$W(0, T, \Omega) \subset L^2(0, T; L^2(\Omega)) \quad (3.2)$$

is compact (Aubin’s Lemma, see [29]).

In the same way, we introduce $W(0, T, \Omega_x)$ the space

$$W(0, T, \Omega_x) := \{\omega \in L^2(0, T; H^1(\Omega_x)), \partial_t \omega \in L^2(0, T; (H^1(\Omega_x))')\},$$

endowed with the Hilbertian norm

$\|\cdot\|_{W(0,T,\Omega_x)}^2 = \|\cdot\|_{L^2(0,T;H^1(\Omega_x))}^2 + \|\partial_t \cdot\|_{L^2(0,T;(H^1(\Omega_x))')}^2$. The same compacity results hold true in this case too.

3.2. Main results. We aim giving an existence result of physically admissible weak solutions for model (\mathcal{M}) completed by initial and boundary conditions.

Let us first detail the mathematical assumptions. We begin with the characteristics of the porous structure. The study is limited to the isotropic case so K_0 is assumed to be a scalar. In the saturated part, the averaged hydraulic conductivity \tilde{K} is thus equal to the constant $\frac{K_0 \rho_0 g}{\mu}$. Without loss of generality, we will assume a zero source term, $Q = 0$. The initial data $P_0 \in H^2(\Omega)$ satisfies the compatibility condition

$$P_0(x, h_0) = P_s \quad \text{in } \Omega_0.$$

Let $\delta \in \mathbb{R}$ be a positive number, we assume that $h_0 \in L^\infty(\Omega_x)$ is such that

$$h_{bot} + \delta \leq h_0 \leq h_{soil} \quad \text{a.e. in } \Omega_x. \quad (3.3)$$

Functions θ and κ are pressure-dependent and we assume

$$\theta \in \mathcal{C}^1(\mathbb{R}), \quad 0 < \theta_- := \phi_{s0} \leq \theta(x) \leq \theta_+, \quad \theta'(x) \geq 0 \quad \forall x \in \mathbb{R}, \quad (3.4)$$

$$\kappa \in \mathcal{C}(\mathbb{R}), \quad 0 < \kappa_- \leq \kappa(x) \leq \kappa_+ \quad \forall x \in \mathbb{R}. \quad (3.5)$$

Before stating the main result of this work, we will transform the original problem and bring us back to the framework introduced in [25].

The above assumptions on the fluid and the medium allow to eliminate the nonlinearity in time of Eq. (2.17), namely assumptions (3.4)-(3.5) are sufficient to define the primitive function \mathcal{P} such that

$$\mathcal{P}(P) = \theta(P) + \alpha_P \int^P \theta(s) ds.$$

A direct computation gives $\mathcal{P}'(P) = \theta'(P) + \alpha_P \theta(P) > \alpha_P \theta_- > 0$, indeed by previous hypothesis, we have $\theta'(P) \geq 0$ and $\theta(P) > \phi s_0$. Since $\theta \in C^1(\mathbb{R})$, there exists $\theta'_+ > 0$ such that $0 \leq \theta' \leq \theta'_+$ on the interval $[P_d, P_s]$.

Since \mathcal{P} is a bijective application, the existence of p such that

$$p = \mathcal{P}(P)$$

is equivalent to the existence of P solution of the Richards problem. The transform \mathcal{P} of Eq. (2.17) is

$$\partial_t p - \frac{1}{\mu} \nabla \cdot \left(\frac{\kappa(\mathcal{P}^{-1}(p))}{(\theta' + \alpha_P \theta)(\mathcal{P}^{-1}(p))} K_0 \nabla p \right) - \frac{\rho_0 g}{\mu} \nabla \cdot (\kappa(\mathcal{P}^{-1}(p)) K_0 \vec{e}_3) = 0.$$

To simplify the presentation, we introduce the notation

$$\tau(p) = \frac{1}{\mu} \frac{\kappa(\mathcal{P}^{-1}(p))}{(\theta' + \alpha_P \theta)(\mathcal{P}^{-1}(p))}.$$

Note that, due to hypotheses (3.4)-(3.5), there exist two positive reals τ_- and τ_+ such that

$$0 < \tau_- := \frac{\kappa_-}{\mu(\alpha_P \theta_+ + \theta'_+)} \leq \tau(p) \leq \tau_+ := \frac{\kappa_+}{\mu \alpha_P \theta_-}. \tag{3.6}$$

Let $l = (h_{soil} - h_{bot})$ be the function (space depending) denoting to the total thickness of the subsoil. We introduce the function T_l defined by

$$T_l(u) = \sqrt{u} + h_{bot} \quad \forall u \in [\delta^2, l^2],$$

which is extended continuously and constantly outside $[\delta^2, l^2]$.

Remark 3.2. *In order to extend the solution p outside the time dependent domain Ω_t , it is necessary to impose on the function h to be less than or equal to a quantity strictly greater than h_{bot} . This is the reason why the small parameter δ is introduced.*

Setting $u = (h - h_{bot})^2$, Eq. (2.24) becomes

$$\frac{S_0}{2} \partial_t u - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u) = - \int_{T_l(u(t,x))}^{h_{soil}(x)} \frac{\partial p}{\partial t} dz. \tag{3.7}$$

Definition: *The definition of the depth h is derived from the construction of u . Namely, for u given by (3.7), we set*

$$h(t, x) := T_l(u). \tag{3.8}$$

Remark 3.3. *This definition of h allows to define the integration domain Ω_t (and thus the interface Γ_t) in the system (3.11)-(3.12). We emphasize that by definition, h always remains in the interval $[h_{bot} + \delta, h_{soil}]$.*

We are led to consider the new problem in (u, p) completed by the boundary and initial conditions:

$$\frac{S_0}{2} \partial_t u - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u) = - \int_{T_l(u(t,x))}^{h_{soil}(x)} \frac{\partial p}{\partial t} dz \quad \text{in } (0, T) \times \Omega_x, \tag{3.9}$$

$$\nabla u \cdot \vec{\nu} = 0 \quad \text{on } (0, T) \times \partial \Omega_x, \quad u(0, x) = (h_0(x) - h_{bot}(x))^2 \quad \text{in } \Omega_x, \tag{3.10}$$

$$\partial_t p - \nabla \cdot (\tau(p) K_0 \nabla p) - \frac{\rho_0 g}{\mu} \nabla \cdot (\kappa(\mathcal{P}^{-1}(p)) K_0 \vec{e}_3) = 0 \quad \text{in } \mathcal{O}_T, \tag{3.11}$$

$$\begin{aligned} p|_{\Gamma_t} &= \mathcal{P}(P_s) \quad \text{in } (0, T), \quad \nabla(\mathcal{P}^{-1}(p) + \rho_0 g z) \cdot \vec{\nu} = 0 \quad \text{on } (0, T) \times (\Gamma_{soil} \cup \Gamma_{ver}), \\ p(0, x, z) &= \mathcal{P}(P_0)(x, z) \quad \text{in } \Omega_0. \end{aligned} \tag{3.12}$$

Remark 3.4. Let $p \in W(0, T; \Omega)$ such that $p = 0$ in \mathcal{O}_T^c . We need to precise the meaning of the term $\int_{h(t,x)}^{h_{soil}(x)} \frac{\partial p}{\partial t} dz$ ($h = T_t(u(t, x))$):

$$\int_{h(t,x)}^{h_{soil}(x)} \frac{\partial p}{\partial t} dz = \int_{h_{bot}(x)}^{h_{soil}(x)} \chi_{z \geq h(t,x)} \frac{\partial p}{\partial t} dz$$

is the function of $(H^1(\Omega_x))'$ such that $\forall v \in H^1(\Omega_x) \subset H^1(\Omega)$, for $\eta_0 > 0$ small enough

$$\begin{aligned} & \left\langle \int_{h(t,x)}^{h_{soil}} \frac{\partial p}{\partial t} dz, v \right\rangle_{H^1(\Omega_x)', H^1(\Omega_x)} \\ &= \left\langle \int_{h_{bot}}^{h_{soil}} \rho_{\eta_0} * \chi_{\{z \geq (h_{bot} + \delta/2)\}} \frac{\partial p}{\partial t} dz, v \right\rangle_{H^1(\Omega_x)', H^1(\Omega_x)} \\ &= \left\langle \frac{\partial p}{\partial t}, \underbrace{\rho_{\eta_0} * \chi_{\{z \geq (h_{bot} + \delta/2)\}}}_{\in V(\Omega)} v \right\rangle_{V'(\Omega), V(\Omega)}, \end{aligned}$$

where $\rho \in C^\infty(\mathbb{R})$, $\rho \geq 0$, with support in the unit ball such that $\int_{\mathbb{R}} \rho(z) dz = 1$. We set $\rho_{\eta_0}(z) = \rho(z/\eta_0)/\eta_0$, η_0 is chosen s.t. $Supp(\rho_{\eta_0} * \chi_{\{z \geq (h_{bot} + \delta/2)\}}) \subset \{z, z \geq (h_{bot} + \delta/4)\}$ and $\rho_{\eta_0} * \chi_{\{z \geq (h_{bot} + \delta/2)\}} = 1$ if $z \geq (h_{bot} + 3\delta/4)$.

The boundary condition satisfied by the unknown p at the interface Γ_t is then reduced to a homogeneous Dirichlet boundary condition. We set $\bar{p} = p - \mathcal{P}(P_s)$. Since $\mathcal{P}(P_s)$ is a constant, the previous system becomes

$$\begin{aligned} \partial_t \bar{p} - \nabla \cdot (\bar{\tau}(\bar{p})) K_0 \nabla \bar{p} - \frac{\rho_0 g}{\mu} \nabla \cdot (\bar{\kappa}(\bar{p}) K_0 \vec{e}_3) &= Q \quad \text{in } \mathcal{O}_T, \\ \bar{p}|_{\Gamma_t} = 0 \text{ in } (0, T), \nabla(\mathcal{P}^{-1}(\bar{p} + \mathcal{P}(P_s)) + \rho_0 g z) \cdot \vec{\nu} &= 0 \text{ on } (0, T) \times (\Gamma_{soil} \cup \Gamma_{ver}), \\ \bar{p}(0, x, z) &= \mathcal{P}(P_0)(x, z) - \mathcal{P}(P_s) \quad \text{in } \Omega_0, \end{aligned}$$

where $\bar{\tau}(\bar{p}) = \tau(\bar{p} + \mathcal{P}(P_s))$ and $\bar{\kappa}(\bar{p}) = \kappa \circ \mathcal{P}^{-1}(\bar{p} + \mathcal{P}(P_s))$. We thus remark, that just renaming functions τ and κ , we come back to the case $\mathcal{P}(P_s) = 0$ on Γ_t . So, from now, we omit the subscript "—" in the previous system and we consider the system (3.11)-(3.12) with $\mathcal{P}(P_s) = 0$.

The method of auxiliary domains introduced in [25] is used. To this end, the function p is extended by zero outside the variable domain Ω_t . So, we consider the following definition of weak solution associated with the system (3.9)-(3.12):

Definition: We call weak solution of problem (3.9)-(3.12), any solution $(u, p) \in W(0, T, \Omega_x) \times W_0(0, T, \Omega)$ s.t. $\forall (\phi_1, \phi_2) \in L^2(0, T; H^1(\Omega_x)) \times L^2(0, T; V(\Omega))$

$$\begin{aligned} & \int_0^T \left(\frac{S_0}{2} \langle \partial_t u, \phi_1 \rangle + \frac{\tilde{K}}{2} \int_{\Omega_x} \nabla' u \cdot \nabla' \phi_1 dx \right. \\ & \left. + \left\langle \frac{\partial p}{\partial t}, \rho_{\eta_0} * \chi_{\{z \geq (h_{bot} + \delta/2)\}} \phi_1 \right\rangle_{V'(\Omega), V(\Omega)} \right) dt = 0, \end{aligned} \tag{3.13}$$

$$\begin{aligned} u(0, x) &= (h_0(x) - h_{bot}(x))^2 \text{ in } \Omega_x, \\ & \int_0^T \left(\langle \partial_t p, \phi_2 \rangle + \int_{\Omega} (\tau(p) \tilde{K}_0 \nabla p + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(p)) \tilde{K}_0 \vec{e}_3) \cdot \nabla \phi_2 d\mathbf{x} \right) dt = 0, \end{aligned} \tag{3.14}$$

$p = 0$ in \mathcal{O}_T^c , $p(0, x, z) = \mathcal{P}(P_0)(x, z)$ in Ω_0 ,

where $\tilde{K}_0 = K_0$ in \mathcal{O}_T and $\tilde{K}_0 = 0$ in \mathcal{O}_T^c .

We state and prove the following existence result:

Theorem 3.5. *Assume that there exist two real numbers θ_- and κ_- such that*

$$\theta(x) \geq \theta_- > 0 \quad \forall x \in \mathbb{R}, \quad \kappa(x) \geq \kappa_- > 0 \quad \forall x \in \mathbb{R}^+.$$

The system (3.9)- (3.12) admits a weak solution (u, p) satisfying

- (a) the function $u \in L^2(0, T; H^1(\Omega_x))$ and $\partial_t u \in L^2(0, T; (H^1(\Omega_x))')$,
 (b) the function $p \in L^2(0, T; V(\Omega))$ and $\partial_t p \in L^2(0, T; V'(\Omega))$.

We can check that the proof of Theorem 3.5 developed below may be easily adapted to the case of a non-zero source term Q .

Proposition 3.6. *Assume that there exists $\delta' > 0$ such that the following a posteriori inequality holds*

$$h_{bot} + \delta' \leq h(t, x) \leq h_{soil} \quad \text{in } (0, T) \times \Omega_x. \quad (3.15)$$

Then, the model (\mathcal{M}) admits a weak solution (P, h) such that

- (a) the function $P \in L^2(0, T; H^1(\Omega))$ and $\partial_t P \in L^2(0, T; (H^1(\Omega))')$;
 (b) the function $h \in L^2(0, T; H^1(\Omega_x))$, $\partial_t h \in L^2(0, T; (H^1(\Omega_x))')$.

The proof of Proposition 3.6 is a direct consequence of Theorem 3.5 as soon as the inequality 3.15 is satisfied. In this case, we turn back to the original problem by considering the inverse transform \mathcal{P}^{-1} .

Nevertheless, as it is done in Proposition 3 of [13], one can introduce sufficiently large “pumping/supply” source terms allowing to control the lower and upper bounds of the solution h (and therefore to guarantee the inequality 3.15). Next section is devoted to the proof of Theorem 3.5.

4. PROOF OF THEOREM 3.5

Let us sketch the global strategy of the proof. The problem is a strongly coupled nonlinear system, so we apply a fixed-point approach to solve it in two steps. First, the system is decoupled and an existence and uniqueness result for each decoupled and linearized problem is established. The decoupled problem in p is solved by introducing a penalized problem and by passing to the limit to come back to the initial linearized problem. Then, we establish compactness results which allow to prove the global existence in time of the initial problem by applying the fixed point Schauder theorem.

4.1. Fixed point step. We now construct the framework to apply the Schauder fixed point theorem (see [16, 34]). For the fixed point strategy, we introduce two convex subsets (W_1, W_2) of $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$, namely

$$\begin{aligned} W_1 &:= \{u \in W(0, T, \Omega_x); u(0) = u_0, \|u\|_{L^2(0, T; H^1(\Omega_x))} \leq C_u \\ &\quad \text{and } \|u\|_{L^2(0, T; (H^1(\Omega_x))')} \leq C'_u\}, \end{aligned}$$

and

$$\begin{aligned} W_2 &:= \{p \in W_0(0, T, \Omega); p(0) = p_0, \|p\|_{L^2(0, T; H^1(\Omega))} \leq C_p \\ &\quad \text{and } \|p\|_{L^2(0, T; (H^1(\Omega))')} \leq C'_p\}, \end{aligned}$$

constants (C_p, C'_p) and (C_u, C'_u) being defined thereafter.

Let $(\bar{u}, \bar{p}) \in W_1 \times W_2$, we begin by considering the unique solution u of the following linearized problem

$$\frac{S_0}{2} \partial_t u - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u) = - \int_{\bar{h}(t,x)}^{h_{soil}(x)} \frac{\partial \bar{p}}{\partial t} dz \quad \text{in } (0, T) \times \Omega_x, \tag{4.1}$$

$$\nabla u \cdot \vec{\nu} = 0 \quad \text{on } (0, T) \times \partial \Omega_x, \quad u(0, x) = (h_0(x) - h_{bot}(x))^2 \quad \text{in } \Omega_x, \tag{4.2}$$

where $\bar{h}(t, x) := T_l(\bar{u}(t, x))$.

Remark 4.1. Note that thanks to the change of variable $u = (h - h_{bot})^2$, Eq. (2.24) (which is nonlinear and degenerate in space and time) has now a parabolic structure.

Lemma 4.2. Let $h_0 \in L^\infty(\Omega_x)$ satisfying (3.3), there exists a unique weak solution $u \in W(0, T, \Omega_x)$ of (4.1)-(4.2) such that

$$\|u\|_{L^2(0,T;H^1(\Omega_x))} \leq C_u \quad \text{and} \quad \|u\|_{L^2(0,T;(H^1(\Omega_x))')} \leq C'_u,$$

where C_u and C'_u only depend on the data of the problem.

Proof. It follows from the classical textbook [21] pp. 178-179 that for every nonnegative function $\bar{u} \in W(0, T, \Omega_x)$ (and \bar{h} s.t. $\bar{h}(t, x) := T_l(\bar{u}(t, x))$) there exists a solution $u \in W(0, T, \Omega_x)$ of the parabolic problem with smooth coefficients

$$\frac{S_0}{2} \partial_t u - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u) = - \int_{\bar{h}(t,x)}^{h_{soil}(x)} \frac{\partial \bar{p}}{\partial t} dz, \tag{4.3}$$

$$\nabla u \cdot \vec{\nu} = 0 \quad \text{on } (0, T) \times \partial \Omega_x, \quad u(0, x) = (h_0 - h_{bot})^2 \quad \text{in } \Omega_x.$$

Multiplying Eq. (4.3) by u and integrating by parts over Ω_x , we obtain

$$\frac{S_0}{2} \frac{d}{dt} \int_{\Omega_x} |u(t, \cdot)|^2 dx + \frac{\tilde{K}}{2} \int_{\Omega_x} |\nabla u|^2 dx \leq \left| \left\langle \int_{\bar{h}(t,x)}^{h_{soil}} \frac{\partial \bar{p}}{\partial t} dz, u \right\rangle_{H^1(\Omega_x)', H^1(\Omega_x)} \right|.$$

Furthermore, by definition of the function T_l , we have

$$\begin{aligned} \left| \left\langle \int_{\bar{h}(t,x)}^{h_{soil}} \frac{\partial \bar{p}}{\partial t} dz, u \right\rangle_{H^1(\Omega_x)', H^1(\Omega_x)} \right| &\leq \|h_{soil} - h_{bot}\|_\infty^{1/2} \|u\|_{H^1(\Omega_x)} \|\partial_t \bar{p}\|_{V'(\Omega)} \\ &\leq \frac{\tilde{K}}{4} (\|u\|_{L^2(\Omega_x)}^2 + \|\nabla u\|_{L^2(\Omega_x)}^2) + \frac{1}{\tilde{K}} \|h_{soil} - h_{bot}\|_\infty \|\partial_t \bar{p}\|_{V'(\Omega)}^2. \end{aligned}$$

Applying Gronwall's inequality in its differential form, we get

$$\|u(t, \cdot)\|_{L^2(\Omega_x)}^2 \leq e^{\frac{\tilde{K}T}{2S_0}} (\|u_0\|_{L^2(\Omega_x)}^2 + \frac{2}{S_0 \tilde{K}} \|h_{soil} - h_{bot}\|_\infty \|\partial_t \bar{p}\|_{L^2(0,T;V'(\Omega))}^2).$$

Combining the previous estimates, we deduce that

$$\|u\|_{L^2(0,T;H^1(\Omega_x))} \leq C(T, S_0, \tilde{K}, \|h_{soil} - h_{bot}\|_\infty, C'_p) := C_u.$$

On the other hand

$$\begin{aligned} \left\| \frac{du}{dt} \right\|_{L^2(0,T;H^1(\Omega_x)')} &= \sup_{\|v\|_{L^2(0,T;H^1(\Omega_x))} \leq 1} \left| \int_0^T \left\langle \frac{du}{dt}, v \right\rangle_{(H^1(\Omega_x))', H^1(\Omega_x)} dt \right| \\ &\leq \frac{2}{S_0} \left(\frac{\tilde{K}}{2} \|u(t, \cdot)\|_{L^2(0,T;H^1(\Omega_x))} + \|h_{soil} - h_{bot}\|_\infty \|\partial_t \bar{p}\|_{L^2(0,T;V'(\Omega))} \right) \end{aligned}$$

$$\leq \frac{2}{S_0} \left(\frac{\tilde{K}}{2} C_u + \|h_{soil} - h_{bot}\|_\infty C'_p \right) := C'_u.$$

The uniqueness of the solution is obvious. Indeed, if u_1 and u_2 are two solutions of (4.1)-(4.2), then $u = u_1 - u_2$ satisfies

$$\begin{aligned} \frac{S_0}{2} \partial_t u - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u) &= 0 \quad \text{in} \quad (0, T) \times \Omega_x, \\ \nabla u \cdot \vec{\nu} &= 0 \quad \text{on} \quad (0, T) \times \partial\Omega_x, \quad u(0, x) = 0 \quad \text{in} \quad \Omega_x. \end{aligned}$$

Following the previous computations, we infer from Gronwall lemma that $u = 0$ a.e. in $(0, T) \times \Omega_x$. This ends the proof of Lemma 4.2. \square

The results stated in the lemma 3.1 require regular non-cylindrical domains in particular with sufficiently regular boundaries (of class C^1 by pieces as mentioned by Mignot). Since in our problem, we can not guarantee as such regularity at the interface h (which is in $W(0, T, \Omega_x)$), a regularization process is used to place our study within the framework of Mignot [25].

We thus regularize h by convolution in space. Let $\psi \in C^\infty(\mathbb{R}^2)$, $\psi \geq 0$, with support in the unit ball such that $\int_{\mathbb{R}^2} \psi(x) dx = 1$. For $\eta > 0$ small enough, we set $\psi_\eta(x) = \psi(x/\eta)/\eta^2$. We extend h by zero outside Ω_x , so we have $h \in C([0, T]; L^2(\mathbb{R}^2)) \cap W(0, T, \mathbb{R}^2)$. We define \tilde{h} by the convolution product with respect to the space variable

$$\tilde{h} = \psi_\eta * h.$$

Its restriction to Ω_x is denoted in the same way. It fulfills $\tilde{h} \in C^\infty(\bar{\Omega}_x)$, and as $\eta \rightarrow 0$, we have

$$\tilde{h} \rightarrow h \text{ strongly in } C([0, T]; L^2(\Omega_x)) \cap L^2(0, T; H^1(\Omega_x)).$$

In Eqs. (3.11)-(3.12), we replace h by \tilde{h} (the substitution appears in the space integration domain Ω_t).

Let $\bar{p} \in W_2$ and $\tilde{h} (= \psi_\eta * h) \in C^\infty(\bar{\Omega}_x)$ where h is given by Lemma 4.2.

We thus consider the following linearized and regularized problem in Ω_T : Find $p_\eta \in W_0(0, T, \Omega)$ s.t. $\forall \phi \in L^2(0, T; V(\Omega))$

$$\begin{aligned} \int_0^T \left(\langle \partial_t p_\eta, \phi \rangle + \int_\Omega (\tau(\bar{p}) \tilde{K}_0 \nabla p_\eta + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) \tilde{K}_0 \vec{e}_3) \cdot \nabla \phi \, dx \right) dt &= 0, \quad (4.4) \\ p_\eta &= 0 \text{ in } \mathcal{O}_T^c \text{ and } p_\eta(0, x, z) = \mathcal{P}(P_0)(x, z) \text{ in } \Omega_0. \quad (4.5) \end{aligned}$$

Proposition 4.3. *For any $\eta > 0$, there exists a unique function p_η in $W_0(0, T, \Omega)$ solution of (4.4)-(4.5). It fulfills the uniform estimates*

$$\|p_\eta\|_{L^2(0, T; H^1(\Omega))} \leq C_p \quad \text{and} \quad \|p_\eta\|_{L^2(0, T; (H^1(\Omega))')} \leq C'_p, \quad (4.6)$$

where C_p and C'_p only depend on the data of the original problem (3.11)- (3.12).

Let us admit for the moment this Proposition whose the proof will be given at the end. From now, we omit the subscript η in p_η (and in u_η).

Let $(\bar{u}, \bar{p}) \in W_1 \times W_2$, Lemma 4.2 and Proposition 4.3 enable to define an application \mathcal{F} such that:

$$\begin{aligned} W(0, T, \Omega_x) \times W_0(0, T, \Omega) &\rightarrow W(0, T, \Omega_x) \times W_0(0, T, \Omega) \\ \mathcal{F}(\bar{u}, \bar{p}) &= (u, p). \quad (4.7) \end{aligned}$$

The end of the present subsection is devoted to the proof of the existence of a fixed point of \mathcal{F} in some appropriate subset. We conclude the proof of Theorem 3.5 by passing to the limit when $\eta \rightarrow 0$.

Lemma 4.4. *The application \mathcal{F} satisfies:*

- *There exists \mathcal{C} a nonempty, closed, convex, bounded set in $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$ satisfying $\mathcal{F}(\mathcal{C}) \subset \mathcal{C}$,*
- *the application \mathcal{F} defined by (4.7) is weakly sequentially continuous in $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$,*
- *there exists $(u, p) \in W_1 \times W_2$ such that $\mathcal{F}(u, p) = (u, p)$.*

Proof. We set $\mathcal{C} = W_1 \times W_2$, the first point of Lemma 4.4 is obvious thanks to Lemma 4.2 and Proposition 4.3. Indeed \mathcal{C} is clearly a nonempty (strongly) closed convex set in $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$.

Regarding the second point of Lemma 4.4, we first note that \mathcal{C} is compact for the weak topology. \mathcal{F} maps $W_1 \times W_2$ into it self. Let now $(v_n)_{n \geq 0} = (\bar{u}_n, \bar{p}_n)_{n \geq 0}$ be any sequence in \mathcal{C} which is weakly convergent in $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$, and let $v = (\bar{u}, \bar{p})$ be its weak limit. We aim to show (as in [26]) that

$$\mathcal{F}(v_n) \rightharpoonup \mathcal{F}(v) \quad \text{in } W(0, T, \Omega_x) \times W_0(0, T, \Omega) \text{ as } n \rightarrow \infty.$$

Since $\mathcal{F}(v_n) \in W_1 \times W_2$ and $W_1 \times W_2$ is weakly compact, it is sufficient to show that there exists a subsequence (v'_n) of (v_n) such that $\mathcal{F}(v'_n) \rightharpoonup \mathcal{F}(v)$. Extracting a subsequence if needed we may assume without loss of generality that $\mathcal{F}(v_n) \rightharpoonup w$ in $W(0, T, \Omega_x) \times W_0(0, T, \Omega)$ as $n \rightarrow \infty$ for some $w = (u, p) \in W_1 \times W_2$, and we have to show that w and $\mathcal{F}(v)$ agree. Set $w_n = \mathcal{F}(v_n)$ ($w_n = (u_n, p_n)$), it follows from Aubin's Lemma that

$$\begin{aligned} w_n &\rightarrow w \text{ in } L^2((0, T) \times \Omega_x) \times L^2((0, T) \times \Omega) \text{ and } w_n(t, x) \rightarrow w(t, x) \text{ a.e.,} \\ v_n &\rightarrow v \text{ in } L^2((0, T) \times \Omega_x) \times L^2((0, T) \times \Omega) \text{ and } v_n(t, x) \rightarrow v(t, x) \text{ a.e.,} \\ \partial_t w_n &\rightharpoonup \partial_t w \text{ in } L^2(0, T; (H^1(\Omega_x))') \times L^2(0, T; V'(\Omega)), \\ \nabla w_n &\rightharpoonup \nabla w \text{ weakly in } L^2((0, T) \times \Omega_x) \times L^2((0, T) \times \Omega). \end{aligned}$$

Thanks to Lebesgue theorem (and the properties of functions τ and T_l) we obtain that $w = \mathcal{F}(v)$ (since $w(0, \cdot) = (u(0, \cdot), p(0, \cdot)) = (u_0, p_0)$ because $w \in \mathcal{C}$) and the proof that $\mathcal{F}|_{\mathcal{C}}$ be weakly sequentially continuous is complete.

It follows from Schauder theorem [34] that there exists $(u, p) \in W_1 \times W_2$ such that $\mathcal{F}(u, p) = (u, p)$. The proof of Lemma 4.4 is thus achieved. \square

We collect the results obtained previously. We can associate with any real number $\eta > 0$ the fixed point $(u_\eta, p_\eta) \in W_1 \times W_2$ of the mapping \mathcal{F} . It is a solution of the system :

$$\partial_t p_\eta - \nabla \cdot (\tau(p_\eta) K_0 \nabla p_\eta) - \frac{\rho_0 g}{\mu} \nabla \cdot (\kappa(\mathcal{P}^{-1}(p_\eta)) K_0 \vec{e}_3) = 0 \quad \text{in } \mathcal{O}_T, \quad (4.8)$$

$$\begin{aligned} p_\eta|_{\Gamma_t} &= \mathcal{P}(P_s) \text{ in } \mathcal{O}_T, \quad \nabla(\mathcal{P}^{-1}(p_\eta) + \rho_0 gz) \cdot \vec{\nu} = 0 \text{ on } (0, T) \times (\Gamma_{soil} \cup \Gamma_{ver}), \\ p_\eta(0, x, z) &= \mathcal{P}(P_0)(x, z) \quad \text{in } \Omega_0. \end{aligned} \quad (4.9)$$

$$\frac{S_0}{2} \partial_t u_\eta - \frac{\tilde{K}}{2} \nabla' \cdot (\nabla' u_\eta) = - \int_{h_\eta(t,x)}^{h_{soil}(x)} \frac{\partial p_\eta}{\partial t} dz \quad \text{in } (0, T) \times \Omega_x, \quad (4.10)$$

$$\nabla u_\eta \cdot \vec{\nu} = 0 \text{ on } (0, T) \times \partial\Omega_x, \quad u_\eta(0, x) = (h_0(x) - h_{bot}(x))^2 \quad \text{in } \Omega_x. \quad (4.11)$$

We can obtain similar estimates for (u_η, p_η) than those derived in Lemma 4.2 and Proposition 4.3. We thus assert the existence of limit functions (extracting a subsequence if needed) $(u, p) \in W(0, T, \Omega_x) \times W_0(0, T, \Omega)$ such that

$$\begin{aligned} (u_\eta, p_\eta) &\rightarrow (u, p) \text{ in } L^2((0, T) \times \Omega_x) \times L^2((0, T) \times \Omega), \\ (u_\eta(t, x), p_\eta(t, x)) &\rightarrow (u(t, x), p(t, x)) \text{ a.e. in } ((0, T) \times \Omega_x) \times ((0, T) \times \Omega), \\ \tilde{h}(t, x) = \psi_\eta * h(t, x) &\rightarrow h(t, x) \text{ a.e. in } (0, T) \times \Omega_x, \\ (\partial_t u_\eta, \partial_t p_\eta) &\rightarrow (\partial_t u, \partial_t p) \text{ in } L^2(0, T; (H^1(\Omega_x))' \times L^2(0, T; V'(\Omega))), \\ (\nabla u_\eta, \nabla p_\eta) &\rightarrow (\nabla u, \nabla p) \text{ weakly in } L^2((0, T) \times \Omega_x) \times L^2((0, T) \times \Omega). \end{aligned}$$

Letting $\eta \rightarrow 0$ in weak formulations resulting from (4.8)-(4.11), we prove the existence of a weak solution (u, p) of problem (3.11)-(3.10). This ends the proof of Theorem 3.5. \square

4.2. Proof of Proposition 4.3. Again, we omit the subscript η in p_η .

The proof of Proposition 4.3 is done by introducing a penalized problem and by passing to the limit to come back to the linearized problem (4.4)-(4.5).

We thus consider the weak solution p of the linearized problem (4.4)-(4.5).

So we look for $p \in W_0(0, T, \Omega)$ such that, $\forall \phi \in L^2(0, T; V(\Omega))$,

$$\int_0^T \left(\langle \partial_t p, \phi \rangle + \int_\Omega (\tau(\bar{p})\tilde{K}_0 \nabla p + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p}))\tilde{K}_0 \vec{e}_3) \cdot \nabla \phi \right) dt = 0.$$

We first remark that the solution of system (4.4)-(4.5) is unique. Indeed, if p_1 and p_2 are two solutions of (4.4)-(4.5), then $q = p_1 - p_2$ satisfies

$$\int_0^T \left(\langle \partial_t q, \phi \rangle + \int_{\Omega_t} \tau(\bar{p})K_0 \nabla q \cdot \nabla \phi \, d\mathbf{x} \right) dt = 0.$$

Then, taking $\phi = q$ and using the fourth point of Lemma 3.1, we conclude that

$$\frac{1}{2} \int_{\Omega_T} q^2(T, x) \, d\mathbf{x} \, dt + \int_0^T \int_{\Omega_t} \tau(\bar{p})K_0 |\nabla q|^2 \, d\mathbf{x} \, dt = 0,$$

since $q(0, \cdot) = 0$. We infer from this equality that $q = 0$ a.e. in $(0, T) \times \Omega$ (since $q = 0$ on the interface Γ_t).

We will define a family of approximate problems which are linear parabolic problems in the cylindrical domain $(0, T) \times \Omega$, and whose the solution restricted to the set \mathcal{O}_T will converge to the solution p of the linearized equation (4.4).

STEP 1. PENALIZED PROBLEMS

Let $\epsilon > 0$, we now consider the following penalized problem on Ω : Find $p_\epsilon \in W_0(0, T, \Omega)$ s.t. $\forall \phi \in L^2(0, T; \mathcal{D}(\bar{\Omega}))$ null in a neighborhood of Γ_{bot}

$$\begin{aligned} &\int_0^T \langle \partial_t p_\epsilon, \phi \rangle \, dt + \int_0^T \int_\Omega (\tilde{K}_0 \tau(\bar{p}) \nabla p_\epsilon + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p}))\tilde{K}_0 \vec{e}_3) \cdot \nabla \phi \, d\mathbf{x} \, dt \\ &+ \int_{\mathcal{O}_T^c} \nabla p_\epsilon \cdot \nabla \phi \, d\mathbf{x} \, dt + \frac{1}{\epsilon} \int_{\mathcal{O}_T^c} p_\epsilon \phi \, d\mathbf{x} \, dt = 0, \end{aligned} \quad (4.12)$$

$$p_\epsilon(0, x, z) = \mathcal{P}(P_0)(x, z) \text{ in } \Omega_0 \text{ and } p_\epsilon(0, x, z) = 0 \text{ in } \Omega \setminus \Omega_0. \quad (4.13)$$

We aim to state that the penalized system (4.12)-(4.13) admits a unique solution p_ϵ which tends to the solution of problem (4.4)-(4.5) when $\epsilon \rightarrow 0$. Eq. (4.12) can

be written, $\forall \phi \in W_0(0, T, \Omega)$

$$\underbrace{\int_0^T \int_{\Omega} \tau(\bar{p}) \tilde{K}_0 \nabla p_{\epsilon} \cdot \nabla \phi \, d\mathbf{x} \, dt + \int_{\mathcal{O}_T^c} \nabla p_{\epsilon} \cdot \nabla \phi \, d\mathbf{x} \, dt + \frac{1}{\epsilon} \int_{\mathcal{O}_T^c} p_{\epsilon} \phi \, d\mathbf{x} \, dt}_{A_{\epsilon}(p_{\epsilon}, \phi)} + \underbrace{\int_0^T \langle \partial_t p_{\epsilon}, \phi \rangle \, dt}_{L_{\epsilon}(\phi)} = - \int_0^T \int_{\Omega} \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) \tilde{K}_0 \vec{e}_3 \cdot \nabla \phi \, d\mathbf{x} \, dt. \tag{4.14}$$

Due to (3.6), we establish that the coefficients of A_{ϵ} are in $L^{\infty}((0, T) \times \Omega)$. Moreover, we have

$$A_{\epsilon}(p, p) \geq \inf(1, K_0 \tau_-, \frac{1}{\epsilon}) \|p\|_{L^2(0, T; V(\Omega))}, \quad \forall p \in L^2(0, T; V(\Omega)).$$

We directly check that L_{ϵ} is a linear form on $L^2(0, T; V(\Omega))$. We thus deduce the existence and uniqueness for the system (4.12)-(4.13).

STEP 2. LIMIT WHEN $\epsilon \rightarrow 0$

We first derive some uniform estimates with respect to ϵ (and also η). Multiplying Eq. (4.14) by p_{ϵ} and integrating by parts over Ω , we thus obtain $\forall s \leq T$

$$\int_0^s \langle \partial_t p_{\epsilon}, p_{\epsilon} \rangle \, dt + \underbrace{\int_0^s \int_{\Omega} \tau(\bar{p}) \tilde{K}_0 |\nabla p_{\epsilon}|^2 \, d\mathbf{x} \, dt + \int_{\mathcal{O}_s^c} |\nabla p_{\epsilon}|^2 \, d\mathbf{x} \, dt + \frac{1}{\epsilon} \int_{\mathcal{O}_s^c} p_{\epsilon}^2 \, d\mathbf{x} \, dt}_{I_1} = - \underbrace{\int_0^s \int_{\Omega} \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) \tilde{K}_0 \vec{e}_3 \cdot \nabla p_{\epsilon} \, d\mathbf{x} \, dt}_{I_2}, \quad \forall \phi \in W_0(0, T, \Omega).$$

Then, applying Lemma 3.1 to the first term, we get

$$|\int_0^s \langle \partial_t p_{\epsilon}, p_{\epsilon} \rangle_{V'(\Omega), V(\Omega)} \, dt| = \frac{1}{2} \left(\int_{\Omega} p_{\epsilon}^2(s, \cdot) \, d\mathbf{x} - \int_{\Omega} p_{\epsilon}^2(0, \cdot) \, d\mathbf{x} \right).$$

Besides

$$|I_1| \geq K_0 \tau_- \|\nabla p_{\epsilon}\|_{L^2(\mathcal{O}_s)}^2 + \|\nabla p_{\epsilon}\|_{L^2(\mathcal{O}_s^c)}^2 + \frac{1}{\epsilon} \int_{\mathcal{O}_s^c} p_{\epsilon}^2 \, d\mathbf{x} \, dt,$$

$$|I_2| \leq \frac{\epsilon_1}{2} \|\nabla p_{\epsilon}\|_{L^2(\mathcal{O}_s)}^2 + \frac{1}{2\epsilon_1} T \text{mes}(\Omega) \left(\frac{\rho_0 g \kappa + K_0}{\mu} \right)^2.$$

By taking $\epsilon_1 = K_0 \tau_-$, we deduce directly from these estimates that the sequence $\{p_{\epsilon}\}$ is bounded in $L^2(0, T; V(\Omega))$ and the sequence $\{\frac{1}{\sqrt{\epsilon}} p_{\epsilon}\}$ is bounded in $L^2(\mathcal{O}_T^c)$.

Indeed, applying Gronwall's Lemma, there exists a constant C_p depending only on the data such that

$$\|p_{\epsilon}\|_{L^2(0, T; V(\Omega))}^2 \leq C_p. \tag{4.15}$$

More precisely, we have

$$\min(1, \frac{K_0 \tau_-}{2}) \|\nabla p_{\epsilon}\|_{L^2((0, T) \times \Omega)}^2 \leq \frac{T}{2 K_0 \tau_-} \text{mes}(\Omega) \left(\frac{\rho_0 g \kappa + K_0}{\mu} \right)^2 + \frac{1}{2} \int_{\Omega} p_0^2 \, d\mathbf{x}.$$

We thus can extract subsequences $\{p_\epsilon\}, \{\frac{1}{\sqrt{\epsilon}}p_\epsilon\}$ (not relabeled for convenience) and there exist $q \in L^2(0, T; V(\Omega))$ and $q' \in L^2(\mathcal{O}_T^c)$ such that

$$p_\epsilon \rightharpoonup q \quad \text{weakly in } L^2((0, T) \times \Omega) \tag{4.16}$$

$$\frac{1}{\sqrt{\epsilon}}p_\epsilon \rightharpoonup q' \quad \text{weakly in } L^2(\mathcal{O}_T^c) \tag{4.17}$$

$$\nabla p_\epsilon \rightharpoonup \nabla p \quad \text{weakly in } L^2((0, T) \times \Omega). \tag{4.18}$$

It results from the first two convergences that

$$\begin{aligned} p_\epsilon|_{\mathcal{O}_T^c} &\rightharpoonup q|_{\mathcal{O}_T^c} \quad \text{weakly in } L^2(\mathcal{O}_T^c) \\ p_\epsilon|_{\mathcal{O}_T^c} &= \sqrt{\epsilon} \times \frac{1}{\sqrt{\epsilon}}p_\epsilon \rightharpoonup 0 \quad \text{weakly in } L^2(\mathcal{O}_T^c), \end{aligned}$$

so

$$q|_{\mathcal{O}_T^c} = 0. \tag{4.19}$$

Moreover, since $q \in L^2(0, T; V(\Omega))$, we infer from the second result of Lemma 3.1, that we can define $\gamma(q)$ on Γ_t . Furthermore thanks to (4.19), $\gamma(q) = 0$ on $\Gamma_t, 0 \leq t \leq T$, and thus $q|_{\mathcal{O}_T} \in F(\mathcal{O}_T)$. We must now check that $q|_{\mathcal{O}_T}$ satisfies (4.4).

It remains to be established that

$$D_t(q|_{\mathcal{O}_T}) \in F'(\mathcal{O}_T) \quad \text{and} \quad D_t(p_\epsilon|_{\mathcal{O}_T}) \rightharpoonup D_t(q|_{\mathcal{O}_T}) \quad \text{in } F'(\mathcal{O}_T).$$

Let $\psi \in B_T(\mathcal{O}_T)$, the function ψ is extended by zero on \mathcal{O}_T^c , the extension denoted $\tilde{\psi}$ thus belongs to $W_0(0, T, \Omega)$. Hence taking $\phi = \tilde{\psi}$ in (4.12) leads to

$$\int_0^T \langle \partial_t p_\epsilon, \tilde{\psi} \rangle dt + \int_0^T \int_\Omega (\tilde{K}_0 \tau(\bar{p}) \nabla p_\epsilon + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) \tilde{K}_0 \vec{e}_3) \cdot \nabla \tilde{\psi} \, d\mathbf{x} \, dt = 0,$$

that we can write as follows

$$\begin{aligned} & - \int_0^T \langle \partial_t \tilde{\psi}, p_\epsilon \rangle dt + \int_0^T \int_\Omega (K_0 \tau(\bar{p}) \nabla p_\epsilon + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) K_0 \vec{e}_3) \cdot \nabla \tilde{\psi} \, d\mathbf{x} \, dt \\ & = \int_{\Omega_0} p_0(\mathbf{x}) \psi_0(0, \mathbf{x}) \, d\mathbf{x}. \end{aligned}$$

By letting $\epsilon \rightarrow 0$, we obtain

$$\begin{aligned} & - \int_0^T \langle \partial_t \psi, q|_{\mathcal{O}_T} \rangle dt + \int_0^T \int_{\Omega_t} (K_0 \tau(\bar{p}) \nabla q|_{\mathcal{O}_T} + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) K_0 \vec{e}_3) \cdot \nabla \psi \, d\mathbf{x} \, dt \\ & = \int_{\Omega_0} p_0(\mathbf{x}) \psi_0(0, \mathbf{x}) \, d\mathbf{x}. \end{aligned}$$

From this equation, we deduce that, taking $\psi \in \mathcal{D}(\bar{\mathcal{O}}_T)$ null in a neighborhood of Γ_t

$$D_t(q|_{\mathcal{O}_T}) - \nabla \cdot (K_0 \tau(\bar{p}) \nabla q|_{\mathcal{O}_T} + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) K_0 \vec{e}_3) = 0.$$

Since $F(\mathcal{O}_T)$ is the closure in $H^{0,1}(\mathcal{O}_T)$ of functions of $\mathcal{D}(\bar{\mathcal{O}}_T)$ null in a neighborhood of Γ_t , the previous equality holds true in $F'(\mathcal{O}_T)$. Moreover the solution $p_\epsilon|_{\mathcal{O}_T}$ of (4.14) verifies

$$D_t(p_\epsilon|_{\mathcal{O}_T}) - \nabla \cdot (K_0 \tau(\bar{p}) \nabla p_\epsilon|_{\mathcal{O}_T} + \frac{\rho_0 g}{\mu} \kappa(\mathcal{P}^{-1}(\bar{p})) K_0 \vec{e}_3) = 0 \quad \text{in } F'(\mathcal{O}_T).$$

Letting $\epsilon \rightarrow 0$ and we infer from convergences (4.16)-(4.18) that

$$\underbrace{\nabla \cdot (\tau(\bar{p})\nabla p_\epsilon|_{\mathcal{O}_T} + \frac{\rho_0 g}{\mu}\kappa(\mathcal{P}^{-1}(\bar{p}))\vec{e}_3)}_{=D_t(p_\epsilon|_{\mathcal{O}_T})} \rightharpoonup \nabla \cdot (\tau(\bar{p})\nabla q|_{\mathcal{O}_T} + \frac{\rho_0 g}{\mu}\kappa(\mathcal{P}^{-1}(\bar{p}))\vec{e}_3),$$

weakly in $F'(\mathcal{O}_T)$. Thus $q|_{\mathcal{O}_T}$ is the unique solution of (4.4)-(4.5), and the limit of $p_\epsilon|_{\mathcal{O}_T}$ being independent of the chosen subsequence, the whole sequence converges towards $q|_{\mathcal{O}_T}$. Moreover, the first inequality in (4.6) is obtained for the solution $q \in L^2(0, T; V(\Omega))$ of system (4.4)-(4.5) in the same way as for the estimate (4.15) obtained for p_ϵ . Finally, as was done in Lemma 4.2, we deduce from the first inequality of (4.6) that

$$\|\partial_t q\|_{L^2(0, T, V'(\Omega))}^2 \leq C'_p,$$

where C'_p depends on the data and on C_p . This ends the proof of Proposition 4.3. \square

REFERENCES

- [1] M. B. Abbott, J. C. Bathurst, J. A. Cunge, P. E. O'connell, J. Rasmussen. An introduction to the european hydrological system - systeme hydrologique europeen,"she", 2: Structure of a physically-based, distributed modelling system. *Journal of Hydrology*, **87(1)** (1986), 61–77.
- [2] P. Ackerer, A. Younes, Efficient approximations for the the simulation of density driven flow in porous media, *Adv. Water Res.*, **Vol. 31** (2008), 15–27.
- [3] H. W. Alt, S. Luckhaus, Quasilinear elliptic-parabolic differential equations, *Math. Z.*, **Vol. 1** (1983), 311–341.
- [4] H. W. Alt, E. Di Benedetto, Nonsteady flow of water and oil through inhomogeneous porous media, *Ann. Scuola Norm. Sup. Pisa*, **12** (1985), 335–392.
- [5] J. Bear, Dynamics of fluids in porous media, Elsevier, New-York (1972).
- [6] J. Bear, Verruijt A., Modeling groundwater flow and pollution, Springer, Netherlands (1987).
- [7] H. Berninger, M. Ohlberger, O. Sander, K. Smetana, Unsaturated subsurface flow with surface water and nonlinear in- and outflow conditions. *Mathematical Models and Methods in Applied Sciences*, **24(05)** (2014), 901–936.
- [8] C. Bourel, C. Choquet, C. Rosier and M. Tsegmid, Modelling of shallow aquifers in interaction with overland water, *Appl. Math. Model.*, **81** (2020), 727–751.
- [9] R.H. Brooks, A.T. Corey, Hydraulic Properties of Porous Media, *Colorado State University Hydrology Papers*, (1964).
- [10] X. Chen, A. Friedman, T Kimura, Nonstationary filtration in partially saturated porous medium, *Euro. J. Appl.*, **5** (1994), 405–429.
- [11] C. Choquet,, M. M. Diédhiou, C. Rosier, Derivation of a Sharp-Diffuse Interfaces Model for Seawater Intrusion in a Free Aquifer. Numerical Simulations, *SIAM J. Appl. Math.*, **76** (2016), no. 1, 138–158.
- [12] C. Choquet,, M. M. Diédhiou, C. Rosier, Mathematical analysis of a sharp-diffuse interfaces model for seawater intrusion, *J. Differential Equations*, **259** (2015), no. 8, 3803–3824.
- [13] C. Choquet, C. Rosier, L. Rosier, Well posedness of general cross-diffusion systems, *J. Differential Equations*, **300** (2021), 386–425.
- [14] H. Darcy, Les fontaines publiques de la ville de Dijon; exposition et application des principes à employer dans les questions de distribution d'eau, Victor Dalmont, Editeur, Paris, (1856).
- [15] J. Dupuit, Etudes théoriques et pratiques sur le mouvement des eaux dans les canaux découverts et à travers les terrains perméables, 2ème édition, Dunod, Paris, (1863).
- [16] L. C. Evans, Partial differential equations, American Mathematical Society (1998).
- [17] C.W. Fetter, Hydrogeology: A shot history, Part 2, *Ground Water*, **42** (2004), 949–953.
- [18] J. Hulshof, N. Wolanski, Monotone flows in N-dimensional partially saturated porous media: Lipschitz continuity of the interface, *Arch. Rat. Mech. Anal.*, **102** (1988), 287–305.
- [19] H. Kong, R.Lian; Free boundary value problem for compressible magnetohydrodynamic equations, *Electron. J. Differential Equations*, **Vol. 2020** (2020), no. 11, 1-16.

- [20] J. Kong, P. Xin, Z. Yao Song, L. Li, A new model for coupling surface and subsurface water flows: With an application to a lagoon, *Journal of Hydrology*, **390(1)** (2010), 116 – 120.
- [21] O.A. Ladyzhenskaja, V. A. Solonnikov and N. N. Ural'ceva, Linear and quasilinear equations of parabolic type, Amer. Math. Soc. (1968).
- [22] J. L. Lions, Sur les problèmes mixtes pour certains systèmes paraboliques dans des ouverts non cylindriques, *Annales de l'Institut Fourier*, **7** (1957), 143-182.
- [23] J. L. Lions, E. Magenes, Problèmes aux limites non homogènes, Vol. 1, Dunod (1968).
- [24] C.M. Marle, Henry Darcy et les écoulements de fluides en milieux poreux, *Oil and Gas science and technology Re. IFP*, **Vol. 61 (5)** (2006), 599-609.
- [25] A. L. Mignot, Méthodes d'approximation des solutions de certains problèmes aux limites linéaires, *Rendiconti del Seminario Matematico della Università di Padova*, **tome 40** (1968), p. 1-138.
- [26] K. Najib, C. Rosier, On the global existence for a degenerate elliptic-parabolic seawater intrusion problem, *Math. Comput. Simulation*, **Vol. 81 Issue 1** (2011), 2282–2295.
- [27] M. F. Pikul, R. L. Street, I. Remson, A numerical model based on coupled one-dimensional richards and boussinesq equations, *Water Resources Research*, **10(2)** (1974), 295–302.
- [28] R.E. Showalter, N Su, Partially saturated flow in a poroelastic medium, *Disc. Cont. Dyn. Syst. Ser. B* (2001), 403-420.
- [29] J. Simon, Compact sets in the space $L^p(0, T, B)$, *Ann. Mat. Pura Appl.*, **vol. 146 (4)** (1987), 65–96.
- [30] P. Sochala, A. Ern, and S. Piperno, Mass conservative bdf-discontinuous galerkin/explicit finite volume schemes for coupling subsurface and overland flows, *Computer Methods in Applied Mechanics and Engineering*, **198(27)** (2009), 2122 – 2136.
- [31] M. Tsegmid, Modélisation d'aquifères peu profonds en interaction avec les eaux de surfaces, Phd Thesis, 2019 ULCO. www.theses.fr/2019DUNK0526
- [32] M. T. Van Genuchten, A closed-form equation for predicting the hydraulic conductivity of unsaturated soils 1, *Soil science society of America journal*, **44(5)** (1980), 892–898.
- [33] H. M. Yin, A singular-degenerate free boundary problem arising from the moisture evaporation in a partially saturated porous medium, *Ann. Mat. Pura Appl.*, **161** (1992), 379-397.
- [34] E. Zeidler, *Nonlinear functional analysis and its applications, Part 1*, Springer Verlag, (1986).

SAFAA AL NAZER

UNIVERSITÉ DU LITTORAL CÔTE D'OPALE, UR 2597, LMPA, LABORATOIRE DE MATHÉMATIQUES PURES ET APPLIQUÉES JOSEPH LIOUVILLE, F-62100 CALAIS, FRANCE

Email address: safaaalnazer9@gmail.com

CAROLE ROSIER

UNIVERSITÉ DU LITTORAL CÔTE D'OPALE, UR 2597, LMPA, LABORATOIRE DE MATHÉMATIQUES PURES ET APPLIQUÉES JOSEPH LIOUVILLE, F-62100 CALAIS, FRANCE

Email address: Carole.Rosier@univ-littoral.fr

MUNKHGEREL TSEG MID

SCHOOL OF APPLIED SCIENCES, MONGOLIAN UNIVERSITY OF SCIENCE AND TECHNOLOGY, ULAANBAATAR, MONGOLIA

Email address: tsemuugii@must.edu.mn